

陕西勉略宁地区铜厂湾金矿床辉绿岩年代学、地球化学特征及其构造意义*

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摘要 勉(县)-略(阳)-宁(强)地区在大地构造位置上处在秦岭造山带与扬子板块西北缘的接合部, 是中国大陆造山系重要的组成部分, 也是重要的矿化集中区, 区内岩浆活动强烈, 对岩浆岩的形成过程进行研究有利于促进对岩浆-成矿活动规律的系统认识。文章通过对勉略宁地区铜厂湾金矿区内辉绿岩脉进行岩石学、锆石 U-Pb 年代学和地球化学研究, 讨论了辉绿岩的成因、物质来源、形成构造背景及其对区域找矿的指示意义。结果表明, 辉绿岩脉的成岩年龄为 $(199.6 \pm 1.8) \text{Ma}$, 为早侏罗世。全岩地球化学特征显示, $w(\text{SiO}_2)$ 为 47.13%~51.56%, $w(\text{MgO})$ 为 3.83%~6.41% ($\text{Mg}^\# = 37 \sim 48$), $w(\text{FeO}^\text{T})$ 为 9.36%~12.74%, $w(\text{TiO}_2)$ 为 1.37%~1.86%, 呈亚碱性, 属拉斑玄武岩系列。岩石稀土元素总量(ΣREE)为 $60.18 \times 10^{-6} \sim 189.06 \times 10^{-6}$, 轻、重稀土元素分馏显著($(\text{La}/\text{Yb})_n = 2.66 \sim 9.12$), 富集轻稀土元素(LREE), 具有右倾型稀土元素配分模式, 具 Eu 的弱负异常($\delta \text{Eu} = 0.6 \sim 0.97$), 与原始地幔相比具有明显的 Nb、Ta、Sr 负异常和 Ba、K 正异常。尽管铜厂湾矿区辉绿岩地球化学组成具有弧岩浆岩的特征, 但根据区域构造演化历史, 笔者认为辉绿岩形成于早侏罗世南秦岭后碰撞阶段, 造山带整体处于减压增温的特殊构造体制, 地幔物质上涌, 底侵作用使得下地壳物质发生部分熔融从而形成了混合型岩浆。

关键词 地球化学; 辉绿岩; 锆石 U-Pb 定年; 铜厂湾金矿床; 勉略宁

中图分类号: P618.51

文献标志码: A

Geochronology and geochemistry of diabase veins from Tongchangwan gold deposit in Mianluening area of Shaanxi Province and its tectonic significance

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Abstract

Located at the juncture of the Qinling orogenic belt and the northwestern margin of the Yangtze plate, the Mianluening triangular zone is an important part of China's orogenic system and mineralization concentration area. This area is characterized by intense magmatic activities, so studies of the formation process of magmatic rocks are helpful to the systematic understanding of magmatic metallogenic activities. In this study, the authors carried out petrologic and geochemical investigation as well as LA-ICP-MS zircon U-Pb dating for diabase in the Tongchangwan gold deposit of the Mianluening triangular area. Accordingly, petrogenesis, material source and

* 本文得到国家重点研发计划(编号:2016YFC0600202)、中国地质科学院基本科研业务费专项经费(编号:K1614)和国家自然科学基金项目(编号:41702213)联合资助

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收稿日期 2020-03-30; 改回日期 2020-10-21。秦思婷编辑。

tectonic setting of diabase and its significance for regional prospecting were discussed. The results show that the diabase was formed at (199.6 ± 1.8) Ma of early Jurassic period. The diabase has SiO_2 content of 47.13%~51.56%, MgO content of 3.83%~6.41% ($\text{Mg}^\# = 37 \sim 48$), FeO^T content of 9.36%~12.74% and TiO_2 content of 1.37%~1.86%, showing subalkaline and tholeiitic features. The diabase shows high total REE content ($60.18 \times 10^{-6} \sim 189.06 \times 10^{-6}$) and $(\text{La}/\text{Yb})_N$ ratios (2.66~9.12) with right-inclined REE patterns, weak negative Eu anomaly ($\delta\text{Eu} = 0.6 \sim 0.97$), depletion of Nb, Ta, Sr and enrichment of Ba and K. Although the geochemical composition of the diabase in the Tongchangwan deposit has the characteristics of arc magmatic rock, the authors, according to the regional tectonic evolution history, believe that the diabase was formed in the post collisional stage of South Qinling in the Early Jurassic period. During the post collisional stage, the orogenic belt was dominated by decompression and heating, accompanied by the upwelling of mantle materials and the partial melting of lower crustal materials, which caused the formation of mixed magma.

Key words: geochemistry, diabase, zircon U-Pb dating, Tongchangwan gold deposit, Mianlueing area

秦岭造山带记录了华北板块和扬子板块之间长期、复杂的地质过程,经历了新元古代、古生代和中生代构造岩浆热事件和造山作用,属于典型的复合型造山带(姜春发,2000;张国伟等,2001;王宗起等,2009a;杨经绥等,2010)。南秦岭构造带作为其中的重要组成部分,对于研究者们了解区域构造演化起着至关重要的作用。虽然,对于南秦岭地区的构造演化模式仍存在较大争议,如认为该地区经历了泥盆纪裂解作用形成有限洋盆以及三叠纪发生向北俯冲而闭合(张国伟等,1995;1996;2003;李曙光等,1996;赖绍聪等,1997;Meng et al., 1999;2000;李三忠等,2002;Li et al., 2004;Lai et al., 2004;Liu et al., 2005;Dong et al., 2011a;2011b),或认为元古代该地区存在原始大洋,经过晚古生代南向俯冲,于三叠纪闭合(王宗起等,1999;2002;2009a;2009b;王涛等,2006;2009;2011;王涛,2008)。但是,研究者均认为造山带自晚三叠世开始进入碰撞阶段,并引发一系列的构造及岩浆活动。在此期间,成矿作用也贯穿于大陆碰撞之后的各阶段,并在不同造山演化阶段发育不同的区域成矿类型,形成了独具特色的矿床组合(侯增谦等,2006)。

基性岩脉是衡量重要构造转换时间的标尺,具有特殊的大陆动力学意义(如陆建军等,2006;柯昌辉等,2020)。深入研究与造山过程相关的基性岩浆岩的时代、类型及成因,对理解造山带构造演化过程及成矿背景具有重要意义(Halls, 1982)。勉(县)-略(阳)-宁(强)构造带地处南秦岭造山带南部,属于秦岭造山带-松潘-甘孜造山带以及扬子板块交接部位,是中国大陆中央造山系重要组成部分(张国伟等,2001)。同时,该地区也是陕西省著名的金三角,

以发育一系列的中小型金多金属矿床为特征,如煎茶岭、东坝沟、铎厂沟、李家沟、铜厂湾等金矿床(孟宪忠,1997;丁振举等,2003a;2003b;张复新等,2004;周振菊等,2011a;2011b;戢兴忠,2016;Liu et al., 2016;丁坤等,2017;王华等,2017),具有很大的找矿潜力,因此,厘清区域构造演化过程也可以指导后续找矿工作。在金属矿床集中分布区域内,基性岩脉十分发育,前人的研究主要着重于探讨基性岩脉与各种金属矿化的关系,对基性岩脉的岩石地球化学特征和成因的研究涉及较少,而对这些岩脉进行研究有助于深入认识区域岩石单元的形成时代和大地构造环境,为建立南秦岭造山带的古构造格局提供有力线索。

因此,本文选取了勉略宁成矿区南侧铜厂湾金矿床内发育的辉绿岩脉作为研究对象,通过系统的岩石学、地球化学分析和LA-ICP-MS锆石U-Pb测年工作,并结合区域上已有的研究成果,阐明基性岩脉成因,探讨其形成时的地球动力学背景,为南秦岭造山带的构造演化提供新证据。

1 区域地质背景

勉略宁构造带是秦岭造山带内重要的构造岩浆活动区(图1a),由NE向勉县-阳平关断裂和NW向勉县-略阳断裂南北夹持,呈向西撒开、向东收敛的三角形“楔状体”(张国伟等,1995)。叶连俊等(1944)将区内出露的岩石单元整体称为碧口群(广义),时代为震旦纪—志留纪。之后,不同学者对其进行了解体(秦克令等,1992;陕西省地质矿产勘查开发局,1995;闫全人等,2002;王训练等,2019),并将其自下

而上划分为大安岩组、碧口群(狭义)和秧田坝岩组(在甘肃省也称横丹群),不同的构造单元均以断裂或韧性剪切带分隔(杨运军等,2017)。大安岩组内部除了酸性火山岩含量较少之外,总体岩性与碧口群类似,两者整体上分布于铜钱坝-枫相院断裂和勉县-阳平关断裂之间(图1b),均由熔岩+火山碎屑岩+正常沉积岩共同构成,其中熔岩表现为细碧岩+角斑岩(中性)+安山岩+石英角斑岩(酸性)+流纹岩组合(王宗起等,2009a)。秧田坝岩组(横丹群)主要由长石砂岩及粉砂质板岩夹变质火山岩组成,以火山岩屑和长石碎屑含量丰富、石英碎屑较少为特点(闫全人等,2004;王鹏,2013),总体分布于铜钱坝-枫相院断裂以北(苗雅娜等,2019)。震旦系南沱组、陡山沱组、灯影组由砂岩、粉砂岩及白云岩组成,呈断裂分隔的断块或透镜体,夹于上述岩石单元中。志留系茂县岩群以粉砂质板岩为主,主体位于勉县-阳平关断裂的南东侧。

区域断裂构造十分发育,北部断裂主要为NW-NWW向,南部则以NEE向为主,另外,区域内还叠加近N-S向断裂束(徐学义等,2002)。由南往北,区域大断裂主要为勉县-阳平关、青木川-关口垭、阳坝-苍社平以及铜钱坝-枫相院断裂(张利亚等,2017)。不同方向的断裂将区内岩石单元截切成许多断块,在不同的断块中地层以单斜或倒转褶皱的样式产出。区域岩浆岩以基性岩为主,包括辉绿岩、辉长辉绿岩及辉长闪长岩等,呈脉状及岩株状在断裂带中产出(图1c),局部出现Cu、Zn矿化。局部还出露少量的灰黑色超基性岩(图1c),出露规模较小,发生了强烈蚀变。

铜厂湾金矿床地处陕西安强县大安镇北侧3 km。构造位置上,矿区位于勉县-阳平关断裂附近,此处断裂走向近NEE,向NNW陡倾。矿区地层为碧口群、陡山沱组和大面积分布的灯影组。碧口群主要由变基性凝灰岩、凝灰质砂岩组成,陡山沱组主要出露第二、三岩段,第二岩段由灰色绢云母粉砂质板岩和绢云母板岩组成,第三岩段由灰黑色含碳粉砂质板岩、灰色薄层微晶灰岩构成,灯影组则主要出露第二岩段,由中厚层-块状白云岩和灰色块状砾屑白云岩以及白云质灰岩组成。矿区岩浆岩以辉绿岩为主,侵位于陡山沱组第三岩段和灯影组第二岩段中,规模较小但数量较多,多以岩枝或岩脉的形式沿构造裂隙充填。

矿区断裂较发育,以NEE向为主,NW向和EW向的次之。矿区内部主要为NEE向F11断裂,横贯矿区,倾向NNW,断裂附近岩石发生了明显的韧性变形且控制着矿体的分布(图2)。NW向和近E-W向断裂分布于NEE断裂附近,一般规模不大。

金矿体主要分布在震旦系灯影组中厚层-块状白云岩、白云质灰岩和侵位其中的辉绿岩脉内部,受NEE向主断裂控制(图2)。矿区内可见沿F11断裂延伸的一条较大的辉绿岩脉与金矿化关系密切(图2)。另外,沿北东方向还发育2条规模较大的钠长岩脉,北东侧钠长岩脉贯入灯影组白云岩中,又被构造破碎带截切,南西侧钠长岩脉则平行于构造破碎带发育,两侧被破碎带围限。旁侧辉绿岩发育钠长石化(图3a、c),在钠长石化辉绿岩与碎裂硅化白云岩接触带中发育金矿化。另外,在与F11斜交的次级断裂中发育较多的小辉绿岩脉,其内部矿化较弱,仅见弱的黄铁矿化。

2 样品与分析方法

本次工作采集了铜厂湾金矿床1260 m中段侵位于白云岩中的未发生钠长石化的辉绿岩,可见明显的港湾状侵入边界。局部可见辉绿岩沿后期的断裂侵位,边界较平直(图3b)。辉绿岩为块状构造,显微镜下观察主要矿物为斜长石和辉石,次要矿物为角闪石和黑云母等,具辉绿结构。斜长石(45%~50%)具聚片双晶,自形-半自形板状,发育绢云母化及高岭土化;辉石(35%~40%)则呈半自形-他形,充填在斜长石的三角空隙内(图3d),多发育绿泥石化;角闪石、黑云母含量较少,零散分布。在反射光下,可见不规则的黄铁矿分布于斜长石的孔隙内。5件辉绿岩样品采自远离矿化区域的、新鲜的不同岩脉,代表了铜厂湾金矿区原始侵位的岩脉。

锆石挑选工作由廊坊市诚信地质服务有限公司完成。室内先将岩石样品粉碎至120目以下,用常规的人工淘洗和电磁选方法富集锆石,然后在双目显微镜下手工逐个精选锆石颗粒。锆石样品的制靶工作和阴极发光(CL)照相由南京宏创地质勘查技术服务有限公司完成。在进行锆石的U-Pb年龄测定前,依据透射光图像和阴极发光图像,样品选择晶形好、无裂隙且无包体的颗粒。

锆石U-Pb定年测试工作在中国地质科学院矿产资源研究所MC-ICP-MS实验室完成。锆石定年

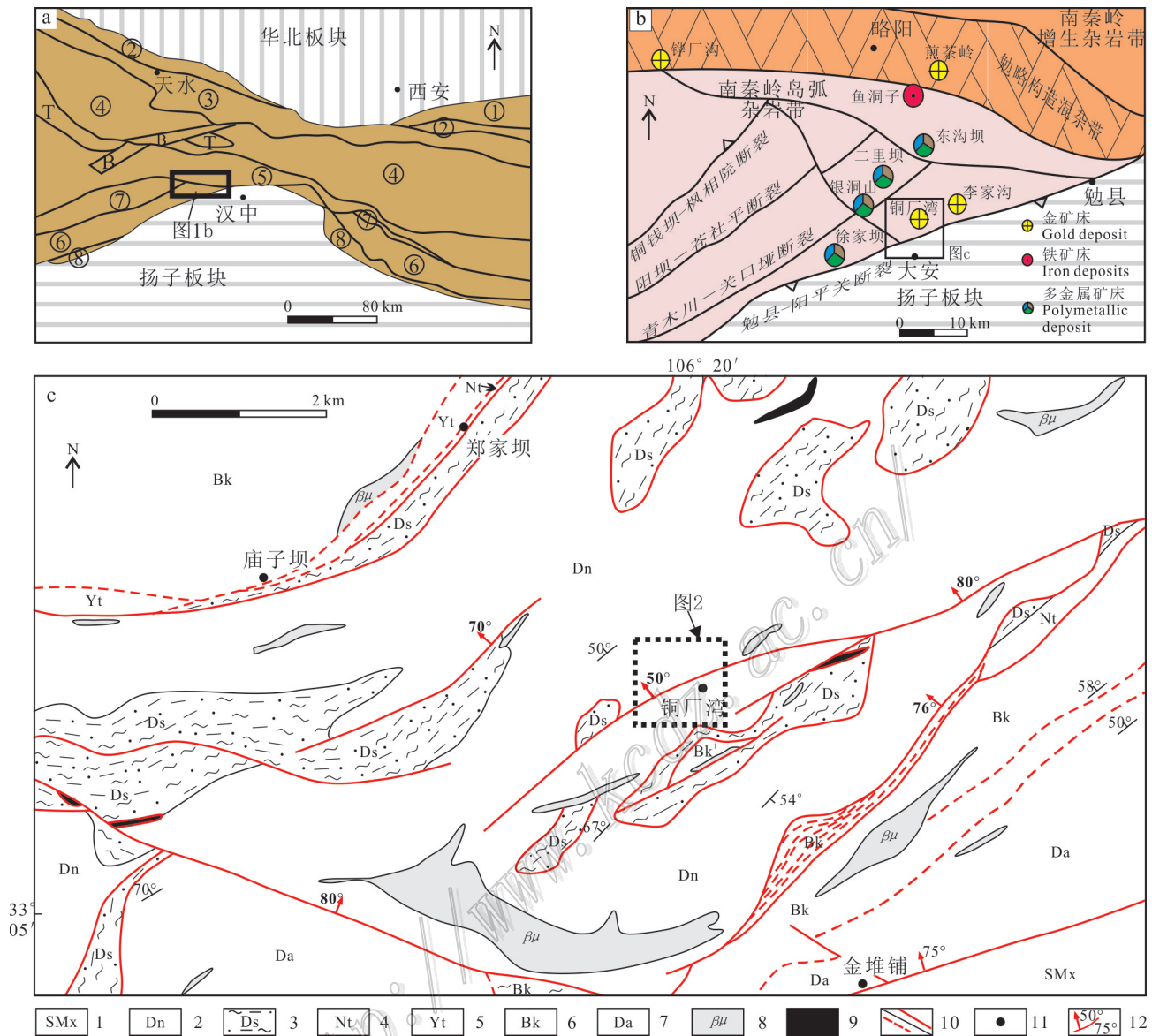


图1 秦岭构造单元划分(a, 据王宗起等, 2009a修改)、勉略宁地区构造简图(b, 据任小华, 2001; 苗雅娜等, 2019修改)

和铜厂湾地区地质图(c, 据陕西省地质矿产勘查开发局, 1995修改)

- ①—华南北缘陆坡带; ②—北秦岭弧后杂岩带; ③—北秦岭岛弧杂岩带; ④—中秦岭弧前盆地系; ⑤—南秦岭增生杂岩带; ⑥—南秦岭岛弧杂岩带; ⑦—南秦岭弧前盆地系; ⑧—南秦岭弧后杂岩带; T—三叠纪残余海盆; B—中生代断陷盆地
1—茂县岩群; 2—灯影组; 3—陡山沱组; 4—南沱组; 5—秧田坝岩组; 6—碧口群; 7—大安岩组; 8—辉绿岩; 9—超基性岩;
10—断裂及韧性剪切带; 11—城镇; 12—产状

Fig. 1 Simplified map of tectonic units of the Qinling orogenic belt (a, modified after Wang et al., 2009a), sketch regional geological map of the Mianluening area (b, modified after Ren, 2001 and Miao et al., 2019) and sketch geological maps of the Tongchangwan area (c, modified after Shaanxi Bureau of Geology and Mineral Exploration and Development, 1995)

- ①—Slope sequence on the southern margin of North China block; ②—Northern back-arc complex; ③—Northern accretionary arc complex;
④—Qinling fore-arc basins; ⑤—Southern accretionary complex; ⑥—Southern arc complex; ⑦—Southern fore-arc basins; ⑧—Southern back-arc complex; T—Triassic remnant sea basins; B—Mesozoic and Cenozoic faulted basins

- 1—Maonian Group; 2—Dingying Formation; 3—Doushantuo Formation; 4—Nantuo Formation; 5—Yangtianba Formation; 6—Bikou Group;
7—Daan Formation; 8—Diabase; 9—Ultrabasic rocks; 10—Fault and ductile shear belt; 11—Town; 12—Attitude

分析所用仪器为Finnigan Neptune型MC-ICP-MS及与之配套的Newwave UP 213激光剥蚀系统。激光

剥蚀孔径为35 μm , 频率为10 Hz, 能量密度约为2.5 J/cm², 以He为载气。锆石U-Pb年龄的测定采用GJ-

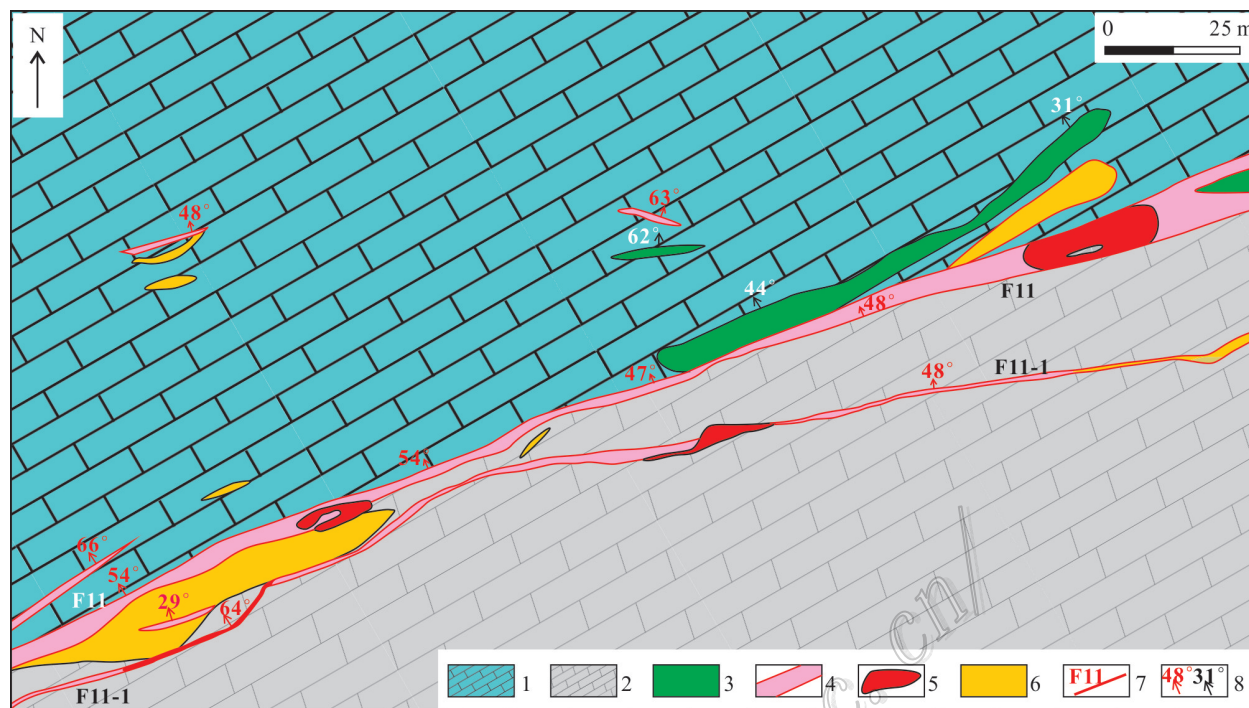


图2 铜厂湾金矿床 1170 m 中段平面图

1—灯影组白云岩;2—灯影组白云质灰岩;3—辉绿岩;4—构造破碎带;5—金矿体;6—钠长岩脉;7—断层及编号;8—产状

Fig. 2 Plan view of mining level 1170 m in the Tongchangwan gold deposit

1—Dengying Formation: dolomite; 2—Dengying Formation: dolomitic limestones; 3—Diabase; 4—Structural fractured zone; 5—Gold orebody; 6—Albite vein; 7—Fault and its serial number; 8—Attitude

1 为外标的校正方法,每隔 10 个样品分析点测 1 次标样,以保证标样和样品的仪器条件完全一致。以 Si 作内标来测定锆石中 U、Pb 和 Th 的含量。数据处理采用 ICPMSData Cal 程序,锆石年龄谱和图用 Iso-plot 3.0 程序完成(Ludwig, 2003)。详细的实验测试过程参考侯可军等(2009)。

硅酸盐全分析在北京燕都中实测试技术有限公司完成,测试流程为:将岩石粉碎粗碎至厘米级的块体,选取无蚀变及脉体穿插的新鲜样品用纯化水冲洗干净,烘干并粉碎至 200 目以备测试使用。主量元素测试首先将粉末样品称量后加 Li₂B₄O₇(1:8)助熔剂混合,并使用融样机加热至 1150℃ 使其在金铂坩埚中熔融成均一玻璃片体,后使用 XRF(Zetium, PANalytical)测试。测试结果保证数据误差小于 1%。微量元素测试将 200 目粉末样品称量并置放入聚四氟乙烯溶样罐并加入 HF+HNO₃,在干燥箱中将高压消解罐保持温度在 190℃ 72 h,后取出经过赶酸并将溶液定容为稀溶液上机测试。测试使用 ICP-MS (M90, analytikjena)完成,所测数据根据监控标样

GSR-2 显示误差小于 5%,部分挥发性元素及极低含量元素的分析误差小于 10%。

3 分析结果

3.1 锆石年龄

辉绿岩中的锆石颗粒大多为无色透明、长柱状,自形-半自形。粒径较小,大部分锆石长度小于 100 μm,长宽比多为 2:1,少量可达 5:1,CL 图像中大部分锆石环带清晰(图 4a)。本次工作对具有震荡环带的 26 颗锆石进行了 U-Pb 同位素测试,分析结果见表 1 和图 4。26 个测点的 Th 和 U 含量分别为 (120~470)×10⁻⁶ 和 (217~735)×10⁻⁶,对应的 Th/U 比值为 0.38~1.13,结合 CL 图像所示岩浆震荡环带结构及晶形等,指示这些锆石为典型的岩浆成因。

26 个年龄数据点均位于一致曲线附近,并组成一个年龄集中区(图 4b),²⁰⁶Pb/²³⁸U 加权平均年龄为 (199.6±1.8) Ma(n=26, MSWD=1.9),代表了辉绿岩的结晶年龄。

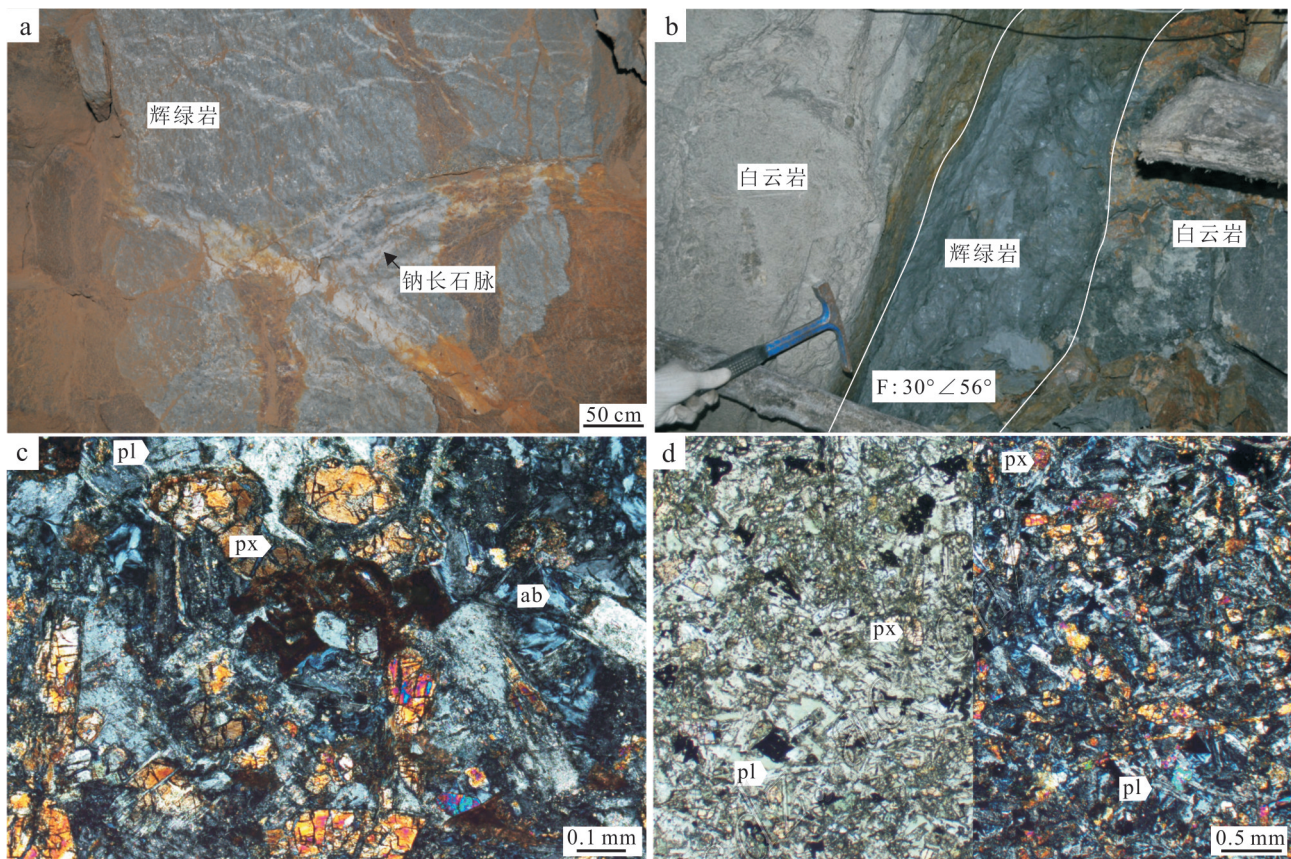


图3 铜厂湾金矿床辉绿岩脉野外和显微照片

a. 辉绿岩内部钠长石脉; b. 辉绿岩脉侵入主断裂旁侧的斜断裂; c. 钠长石化辉绿岩; d. 辉绿结构(左为单偏光、右为正交光)
ab—钠长石; px—辉石; pl—斜长石

Fig. 3 Outcrops and microphotographs of diabases in the Tongchangwan gold deposit

a. Albite veins in the diabase; b. Diabase intrusion along the oblique faults near the major fault; c. Albitized diabase;
d. Diabasic texture (left, plane polarized light; right, cross polarized light)
ab—Albite; px—Pyroxene; pl—Plagioclase

3.2 主量、微量元素

辉绿岩的主微量测试结果见表2,岩石的 $w(\text{SiO}_2)$ 为47.13%~51.56%、 $w(\text{MgO})$ 为3.83%~6.41%($\text{Mg}^\# = 37\sim 48$)、 $w(\text{TiO}_2)$ 为1.37%~1.86%、 $w(\text{FeO}^\text{T})$ 为9.36%~12.74%,其他主量元素 $w(\text{Na}_2\text{O})$ 为0.07%~3.69%、 $w(\text{K}_2\text{O})$ 为0.31%~1.71%、 $w(\text{CaO})$ 为5.19%~10.07%、 $w(\text{Al}_2\text{O}_3)$ 为14.63%~16.39%和 $w(\text{P}_2\text{O}_5)$ 为0.17%~0.33%,全碱含量低($w(\text{Na}_2\text{O} + \text{K}_2\text{O}) = 0.38\%\sim 5.40\%$)。尽管其中的2个样品(17HC38-5/-7)烧失量相对较高(图5中带黑色边框的2个数据点),其主量元素判别存在一定偏差,但在 $\text{Zr}/\text{TiO}_2 \times 0.0001$ - SiO_2 图解(图5a)中,辉绿岩样品大多投在亚碱性系列区域内,在AFM图解(图5b)中,样品均投到拉斑玄武岩系列

区,表明辉绿岩属于拉斑玄武岩系列。

辉绿岩样品稀土元素总量(ΣREE)变化较大($60.18 \times 10^{-6} \sim 189.06 \times 10^{-6}$),轻、重稀土元素比值(LREE/HREE)为3.47~7.42(表2)。在球粒陨石标准化稀土元素配分曲线图(图6a)中,呈右倾配分模式, $(\text{La}/\text{Yb})_\text{N} = 2.66\sim 9.12$, $\delta\text{Eu} = 0.6\sim 0.97$,具有明显的轻、重稀土元素分馏特征和微弱负Eu异常,反映在岩浆演化过程中无明显的斜长石分离结晶作用。总体看来,样品配分曲线位于岛弧拉斑玄武岩和洋岛玄武岩配分曲线之间(图6a)。在原始地幔标准化微量元素图(图6b)中,样品相对亏损Nb、Ta等元素,富集Ba、K,具有与岛弧拉斑玄武岩相一致的元素组成特征。

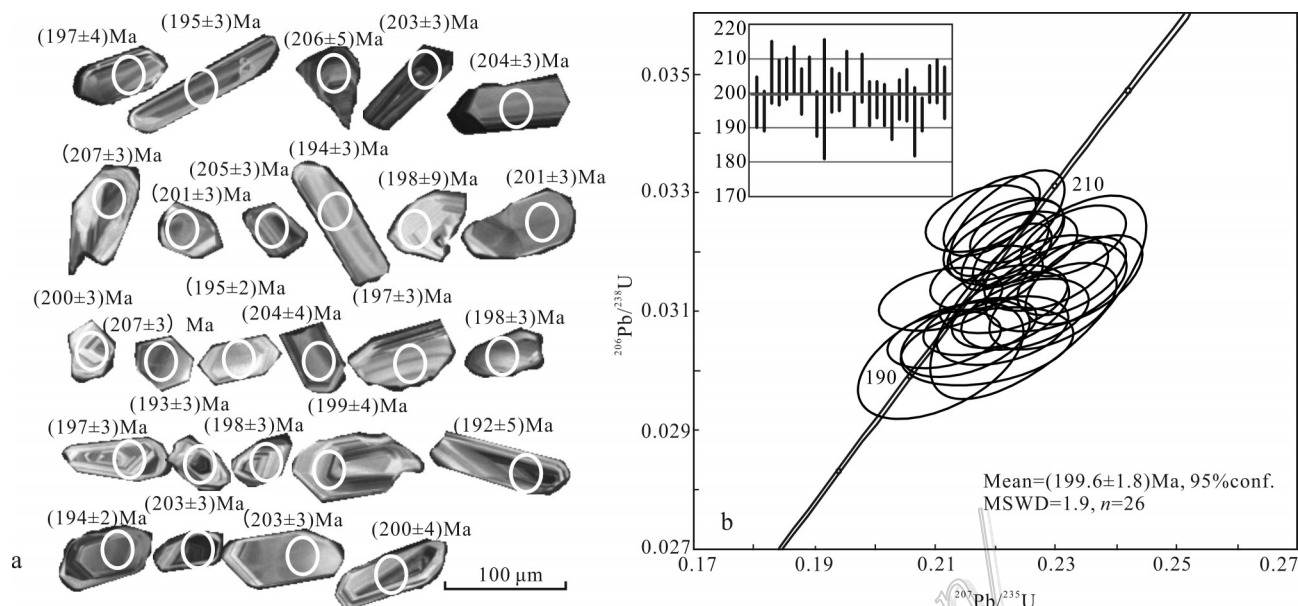


图4 铜厂湾金矿床辉绿岩锆石CL图像、测点位置(a)及年龄谱和图(b)

Fig. 4 Zircon cathodoluminescence images, analytical spots (a) and concordia diagrams (b) of diabase in the Tongchangwan gold deposit

4 讨论

4.1 辉绿岩形成时代

对铜厂湾金矿区内发育的辉绿岩脉进行锆石LA-ICP-MS U-Pb定年,获得 (199.6 ± 1.8) Ma加权平均年龄值,表明辉绿岩均形成于早侏罗世。辉绿岩脉直接侵位于白云岩中或沿白云岩内部的脆性断裂侵位,野外观察到辉绿岩与白云岩之间接触面为港湾状或近平直,表明辉绿岩的侵位与脆性断裂(控矿构造F11及其分支断裂)同时或稍晚。区域上,铜厂湾矿区位于勉县-阳平关断裂附近,F11断裂与阳平关断裂小角度相交,两者属于统一的断裂系统。阳平关断裂作为碧口地块的边界断裂,形成于碧口地块与周边地块相互作用过程中。虽然对于碧口地块的成因存在不同认识,如认为它是从秦岭“峰腰”挤出的刚性构造外来体(王二七等,2001),或认为它是从共和附近三叉裂谷向东运移过来的块体(李三忠等,2002)。但是,均认为在晚三叠世碧口地块开始与秦岭造山带接触,并形成勉县-阳平关断裂。

前人对勉县-阳平关断裂进行了大量的构造变形研究,取得了许多重要进展和认识:文德华(1994)和王二七等(2001)指出在印支期,断裂主要受到NW-SE方向的挤压作用,并伴有左行走滑运动。李

三忠等(2002)认为,断裂记录了印支期间的韧性逆冲(D1)、左行脆韧性斜冲(D2)和右行脆性走滑(D3)3期构造形迹。任清军等(2011)将断裂的变形期次分为六期,其活动时限为印支晚期到喜马拉雅期(表3)。这些研究指出,勉县-阳平关断裂及其分支断裂的脆性运动时限为印支晚期。本次研究的辉绿岩,从年龄到野外地质特征等方面均与前人的认识相一致,证明区域上存在一期早侏罗世的构造事件,即由韧性向脆性构造发生转换,也指示区域上挤压作用的逐渐变弱,以走滑断层和脆性断层为代表的伸展构造开始调节造山带正向的缩短。

4.2 岩石成因

铜厂湾矿区辉绿岩表现为亚碱性、拉斑玄武岩特征,拉斑系列玄武岩可形成于多种构造环境,如洋岛、洋中脊、岛弧、边缘海盆、活动大陆边缘和稳定大陆。辉绿岩样品的 $w(\text{SiO}_2)$ 为47.13%~51.56%、 $w(\text{TiO}_2)$ 为1.37%~1.86%以及 FeO^T/MgO 比值(1.93~3.08),指示它们与岛弧环境下的拉斑玄武岩具有相似的特征(Miyashiro, 1975)。在2Nb-Zr/4-Y图解和Hf/3-Th-Ta图解(图7a、b)中,样品点分别投在板内拉斑及火山弧玄武岩区和岛弧玄武岩区;在Ta/Hf-Th/Hf图解和Ta/Yb-Th/Yb图解(图7c、d)中,样品点分别投在大陆边缘岛弧区和活动大陆边缘区。稀土和微量元素曲线图中,辉绿岩表现为轻稀土元素

表 1 铜厂湾金矿床内辉绿岩锆石的 U-Th-Pb 同位素分析结果

Table 1 U-Th-Pb isotopic analyses for zircon from diabase in the Tongchangwan gold deposit

点号	$w(\text{B})/10^{-6}$		Th/U	同位素比值						年龄/Ma					
	Th ²³²	U ²³⁸		²⁰⁷ Pb/ ²⁰⁶ Pb	1 σ	²⁰⁷ Pb/ ²³⁵ U	1 σ	²⁰⁶ Pb/ ²³⁸ U	1 σ	²⁰⁷ Pb/ ²⁰⁶ Pb	1 σ	²⁰⁷ Pb/ ²³⁵ U	1 σ	²⁰⁶ Pb/ ²³⁸ U	1 σ
1	242	307	0.79	0.0536	0.0019	0.2295	0.0086	0.0311	0.0006	354	78	210	7	197	4
2	208	357	0.58	0.0521	0.0020	0.2211	0.0089	0.0307	0.0005	300	82	203	7	195	3
3	130	343	0.38	0.0499	0.0016	0.2219	0.0078	0.0325	0.0007	187	69	203	6	206	5
4	301	735	0.41	0.0498	0.0010	0.2216	0.0059	0.0320	0.0005	187	46	203	5	203	3
5	120	312	0.39	0.0528	0.0016	0.2328	0.0071	0.0322	0.0005	317	66	213	6	204	3
6	232	422	0.55	0.0487	0.0013	0.2193	0.0065	0.0326	0.0005	132	56	201	5	207	3
7	253	367	0.69	0.0532	0.0015	0.2326	0.0074	0.0316	0.0005	339	95	212	6	201	3
8	248	471	0.53	0.0500	0.0013	0.2235	0.0063	0.0323	0.0004	195	61	205	5	205	3
9	453	401	1.13	0.0512	0.0019	0.2150	0.0087	0.0305	0.0005	256	82	198	7	194	3
10	148	256	0.58	0.0520	0.0022	0.2271	0.0146	0.0312	0.0014	287	101	208	12	198	9
11	218	540	0.40	0.0521	0.0017	0.2284	0.0084	0.0316	0.0005	300	76	209	7	201	3
12	172	296	0.58	0.0504	0.0018	0.2185	0.0076	0.0316	0.0004	213	83	201	6	200	3
13	156	293	0.53	0.0485	0.0016	0.2175	0.0075	0.0326	0.0005	124	78	200	6	207	3
14	207	312	0.66	0.0510	0.0016	0.2174	0.0077	0.0307	0.0004	239	72	200	6	195	2
15	176	397	0.44	0.0495	0.0015	0.2213	0.0082	0.0322	0.0006	172	70	203	7	204	4
16	470	476	0.99	0.0521	0.0017	0.2224	0.0080	0.0310	0.0005	287	74	204	7	197	3
17	225	289	0.78	0.0492	0.0020	0.2109	0.0084	0.0312	0.0004	167	98	194	7	198	3
18	163	251	0.65	0.0528	0.0020	0.2248	0.0086	0.0310	0.0005	320	118	206	7	197	3
19	243	294	0.83	0.0517	0.0025	0.2186	0.0117	0.0304	0.0005	272	111	201	10	193	3
20	273	271	1.01	0.0514	0.0019	0.2213	0.0087	0.0312	0.0005	261	83	203	7	198	3
21	214	225	0.95	0.0531	0.0020	0.2300	0.0095	0.0314	0.0006	345	81	210	8	199	4
22	234	335	0.70	0.0508	0.0026	0.2109	0.0112	0.0302	0.0008	232	119	194	9	192	5
23	283	440	0.64	0.0503	0.0018	0.2115	0.0077	0.0305	0.0004	206	81	195	6	194	2
24	458	602	0.76	0.0508	0.0012	0.2227	0.0053	0.0319	0.0004	232	56	204	4	203	3
25	214	217	0.99	0.0503	0.0018	0.2228	0.0088	0.0321	0.0005	209	90	204	7	203	3
26	255	412	0.62	0.0537	0.0017	0.2346	0.0079	0.0315	0.0006	367	70	214	7	200	4

(LREE)和大离子亲石元素(LILEs: K、Ba、Th、Pb)富集,高场强元素(HFSEs: Nb、Ta、Ti)亏损,其配分曲线与大多数板内玄武岩(MORB、OIB等)存在明显区别,而与大洋板块俯冲有关的岛弧拉斑玄武岩的配分型式类似(图 6a、b)。另外,微量元素配分图中明显的负 Nb、Ta 异常,表明岩石中存在大量大陆物质的参与(Crawford et al., 1987; Davidson, 1987; Jahn et al., 1999; Martin, 1999)。

地壳物质的加入往往造成岩石中大离子亲石元素和轻稀土元素富集,以及 Th、U、Zr、Hf 等元素的正异常(Taylor et al., 1985; Rudnick et al., 2003)。地壳

物质可以通过俯冲进入岩石圈地幔,以熔融产生的熔体形式对岩石圈地幔进行交代改造,也可以在幔源岩浆上升侵位过程中通过地壳混染作用加入。辉绿岩的 Ba/Nb、La/Nb 比值分别为 13.04~84.39、1.57~4.94,也明显高于 MORB、OIB、碱性玄武岩、金伯利岩的相应比值(1~20、0.5~2.5),这表明大陆物质在辉绿岩岩浆中占有重要地位(Dungan et al., 1986)。由于地壳的 La/Nb 比值(2.2)比样品中的比值仍较低,表明铜厂湾矿区辉绿岩呈现的地球化学特征是由源区交代作用引起的。具有弱 Eu 负异常($\delta\text{Eu}=0.6\sim0.97$)的特征,表明无明显的斜长石分离结晶作用,样

表 2 铜厂湾金矿床内辉绿岩的主量元素和微量元素分析结果

Table 2 Major and trace element content of diabase in the Tongchangwan gold deposit

组分	17HC38-3	17HC38-4	17HC38-5	17HC38-6	17HC38-7	组分	17HC38-3	17HC38-4	17HC38-5	17HC38-6	17HC38-7
$w(B)/\%$						$w(B)/10^{-6}$					
SiO ₂	51.24	50.03	47.13	47.92	51.56	Ba	400.98	647.10	242.29	303.63	85.43
TiO ₂	1.86	1.67	1.72	1.37	1.73	La	16.66	37.91	14.86	10.04	10.31
Al ₂ O ₃	15.07	16.39	14.63	14.78	14.92	Ce	34.54	71.40	32.27	21.35	21.91
Fe ₂ O ₃	4.96	2.23	2.61	4.18	1.91	Pr	4.81	9.54	4.56	3.24	3.14
FeO	7.51	7.35	10.39	7.49	9.59	Nd	19.84	37.92	18.57	14.03	12.53
MnO	0.16	0.12	0.12	0.17	0.10	Sm	3.96	7.64	3.64	3.38	2.43
MgO	3.89	3.83	6.41	5.82	4.65	Eu	1.24	2.18	0.92	1.00	0.47
CaO	7.07	7.24	5.19	10.07	5.30	Gd	3.74	6.81	3.15	3.32	2.28
Na ₂ O	3.24	3.69	0.88	3.69	0.07	Tb	0.68	1.02	0.57	0.66	0.39
K ₂ O	0.78	1.71	0.81	1.01	0.31	Dy	4.38	6.53	3.63	4.43	2.51
P ₂ O ₅	0.26	0.33	0.28	0.17	0.28	Ho	0.97	1.30	0.81	1.01	0.59
烧失量	2.96	4.26	8.54	2.48	8.29	Er	2.38	3.06	2.03	2.53	1.50
总和	99.00	98.85	98.71	99.15	98.71	Tm	0.37	0.48	0.34	0.41	0.26
FeO ^T	11.97	9.36	12.74	11.25	11.31	Yb	2.21	2.80	2.18	2.54	1.53
Mg [#]	37	42	47	48	42	Lu	0.36	0.46	0.35	0.40	0.31
FeO ^T /MgO	3.08	2.44	1.99	1.93	2.43	Hf	3.14	3.98	2.89	2.63	2.97
$w(B)/10^{-6}$						Ta	0.29	0.35	0.30	0.20	0.30
Li	6.97	7.66	13.84	6.50	14.46	Pb	2.20	0.81	2.81	3.21	2.09
Sc	34.76	34.96	31.76	40.20	31.88	Th	2.35	4.14	2.36	1.39	2.39
V	432.15	306.43	394.55	336.69	394.86	U	0.57	0.86	1.01	0.38	0.62
Ni	9.80	11.41	17.93	14.93	7.94	ΣREE	96.14	189.06	87.87	68.33	60.18
Co	31.20	33.91	33.14	43.41	34.14	LREE	81.05	166.59	74.82	53.04	50.80
Cu	24.87	33.60	16.72	107.72	31.31	HREE	15.09	22.46	13.05	15.29	9.38
Ga	21.08	21.37	20.99	19.04	21.02	LREE/HREE	5.37	7.42	5.73	3.47	5.42
Rb	14.15	29.72	15.30	13.81	5.87	(La/Yb) _N	5.09	9.12	4.59	2.66	4.54
Sr	312.94	210.47	50.18	213.77	27.17	(La/Sm) _N	2.65	3.12	2.57	1.87	2.67
Y	27.92	36.65	20.23	29.18	15.60	(Gd/Yb) _N	1.37	1.96	1.17	1.05	1.20
Zr	122.56	162.29	115.01	106.52	121.90	δEu	0.97	0.91	0.81	0.90	0.60
Nb	6.37	7.67	6.45	4.42	6.55	δCe	0.92	0.88	0.94	0.90	0.92
Cs	1.76	1.43	1.46	0.73	0.40						

注:比值单位为1;Mg[#]=MgO/(MgO+FeO^T),为分子比。

品固结指数(SI=32.85~67.06)相对较高,也说明其结晶分离程度低。综上所述,铜厂湾矿区辉绿岩全岩地球化学特征表现出类似岛弧玄武岩的特征,岩浆源区受到了壳源物质对岩石圈地幔的交代。

4.3 辉绿岩的构造意义

Dong 等(2011a)对南北秦岭构造带的演化阶段

进行了详细划分,包括北秦岭前寒武纪和早古生代的演化过程、晚古生代—中生代南秦岭的俯冲碰撞过程以及侏罗纪—白垩纪陆内造山过程;刘树文等(2012)通过对区域岩浆岩进行研究将秦岭地区中生代划分为勉略洋盆闭合(248~216 Ma)、北秦岭岛弧杂岩碰撞-造山带垮塌(215~201 Ma)和后碰撞拆沉作

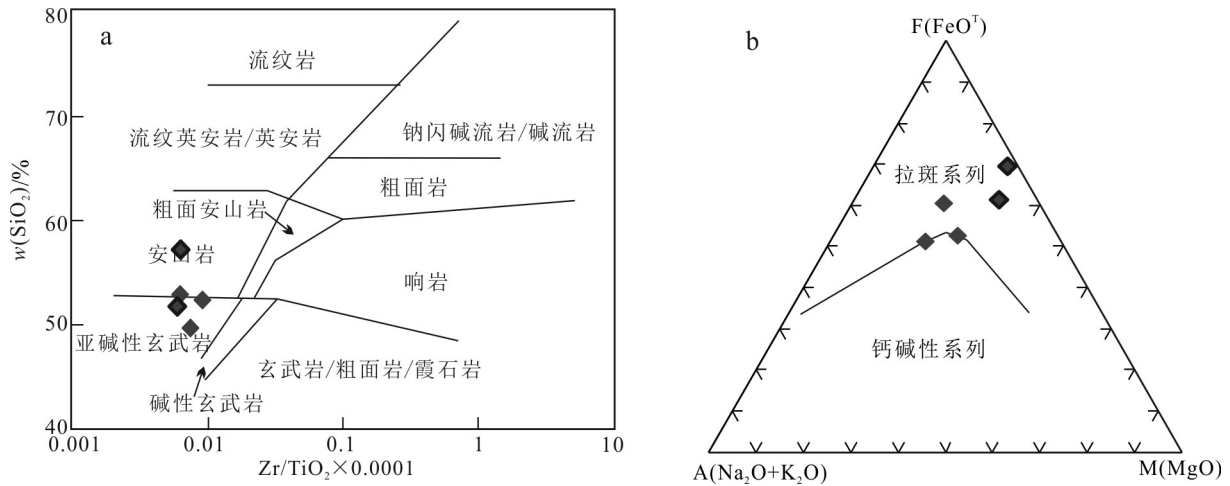


图5 铜厂湾金矿床辉绿岩 $Zr/TiO_2 \times 0.0001-SiO_2$ (a, 据 Winchester et al., 1977) 和 AFM ($A=Na_2O+K_2O$, $F=Total\ FeO$, $M=MgO$) (b, 据 Irving et al., 1971) 图解

Fig. 5 $Zr/TiO_2 \times 0.0001-SiO_2$ plot (a, after Winchester et al., 1977) and AFM ($A=Na_2O+K_2O$, $F=Total\ FeO$, $M=MgO$) plot (b, after Irving et al., 1971) of diabases in the Tongchangwan gold deposit

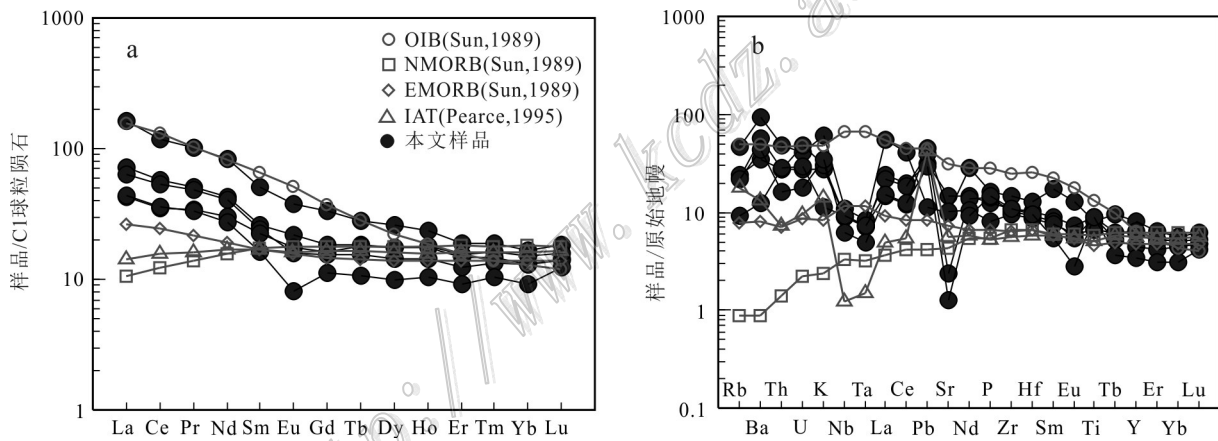


图6 铜厂湾金矿床辉绿岩的球粒陨石标准化稀土元素配分曲线(a)和原始地幔标准化微量元素蛛网图(b)(CI球粒陨石和原始地幔值引自 Sun et al., 1989)

Fig. 6 Chondrite-normalized REE patterns (a) and primitive mantle-normalized spider diagrams (b) of diabases in the Tongchangwan gold deposit (CI chondrite and primitive mantle values after Sun et al., 1989)

表 3 勉县-阳平关断裂变形阶段划分

Table 3 Regional deformation stages of the Mianxian-Yangpingguan fault

对比项目	早-中三叠世 (250~240 Ma)	晚三叠世早期 (240~220 Ma)	晚三叠世—早白垩世 (220~150 Ma)	白垩纪 (150 Ma—新生代)	数据来源
变形阶段	D ₁ NE-SW 挤压	D ₂ NW-SE 挤压	D ₃ 南北挤压	D ₄₋₆ 脆性变形	任清军等, 2011
	D ₁ NW-SE 挤压, 韧性逆冲推覆		D ₂ 左行韧脆性走滑	D ₃ 右行脆性走滑	李三忠等, 2002
		NW-SE 挤压, 伴生左行走滑			曾佐勋等, 2001
		NW-SE 挤压			文德华, 1994
			NNW-SSE 挤压, 兼具左行走滑		王二七等, 2001

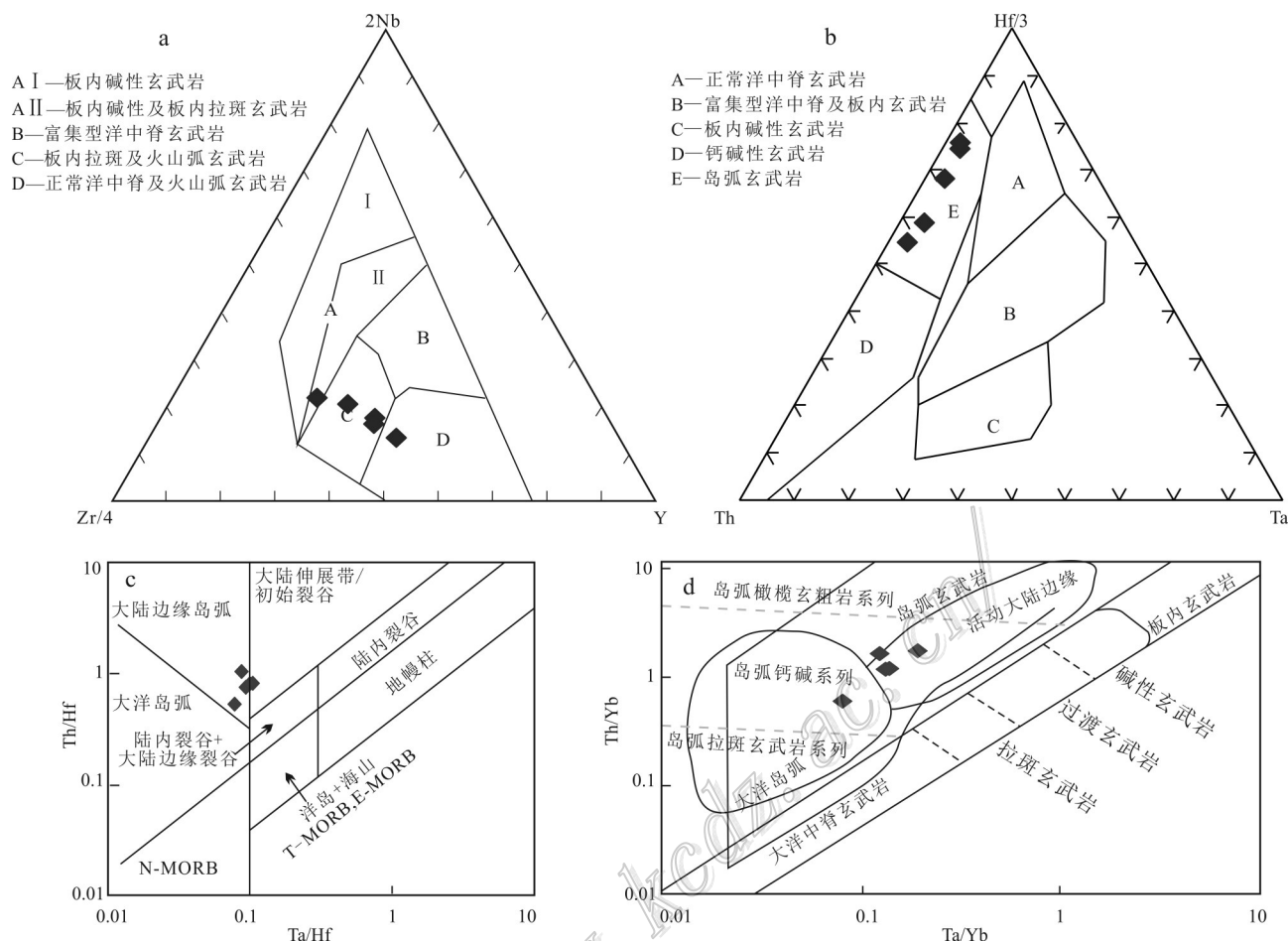


图7 铜厂湾金矿床辉绿岩构造环境判别图

a. 2Nb-Zr/4-Y 图解(底图据 Meschede, 1986); b. Hf/3-Th-Ta 图解(底图据 Wood, 1980); c. Th/Hf-Ta/Hf 图解(底图据汪云亮等, 2001); d. Ta/Yb-Th/Yb 图解(底图据 Pearce, 2008)

Fig. 7 Tectonic setting discrimination diagrams of diorites in the Tongchangwan gold deposit

a. 2Nb-Zr/4-Y (after Meschede, 1986); b. Hf/3-Th-Ta (after Wood, 1980); c. Th/Hf-Ta/Hf (after Wang et al., 2001); d. Ta/Yb-Th/Yb (after Pearce, 2008)

用(200~195 Ma)等阶段;Chen等(2020)通过总结区域上的沉积、变质变形以及岩浆岩数据,也提出西秦岭经历了俯冲(248~235 Ma)、同碰撞(235~225 Ma)以及后碰撞(225~195 Ma)的阶段。研究表明,约200 Ma时期,秦岭造山带已进入后碰撞阶段,这一地质过程也得到其他学者从岩浆岩(张成立等, 2008; 吴峰辉等, 2009; 陈旭等, 2009; Jiang et al., 2010; Qin et al., 2010; Wang et al., 2011; Dong et al., 2012; 吕星球等, 2014; 郑俊等, 2015; Zhang et al., 2019)、变质岩(张宗清等, 2002; Wan et al., 2005; Liu et al., 2006)和构造变形(Li et al., 1999)等方面的证据支持。本文获得碧口铜厂湾地区辉绿岩的早侏罗世年龄为(199.6±1.8)Ma,其结晶年龄与区域上后碰撞阶段的时限一致,这也限定了辉绿岩的形成背景。在同一构造带

上,孙延贵等(2004)也获得了西秦岭-东昆仑结合部基性岩墙群早侏罗世的氩氩坪年龄。这些成果表明,早侏罗世南秦岭勉略洋盆封闭以及后碰撞阶段存在一期构造-岩浆事件。

一个完整的碰撞造山事件应包括挤压、挤压向伸展转变和伸展3个阶段。其中,在挤压向伸展过渡阶段,造山带整体处于减压增温的特殊构造体制,有利于熔融作用的发生,从而使得此阶段会发生明显的壳幔相互作用及岩浆活动(Zhu et al., 2009)。早侏罗世秦岭造山带应力状态发生变化,处于由挤压向伸展过渡阶段(冯益民等, 2003),造山带发生减压,秦岭地区地幔物质上涌,底侵作用使得下地壳物质发生部分熔融从而形成了混合型岩浆,同时,在浅部形成一系列的张性构造,岩浆沿着张性裂隙等软

弱带上升侵位,从而形成了铜厂湾矿区内的辉绿岩脉,它们代表了拉张环境侵位的基性岩脉。

虽然地球化学数据表明,铜厂湾金矿床内部辉绿岩具有大陆边缘弧玄武岩的特征,然而,也有一些研究指出,具火山弧特征的岩浆岩并不一定指示俯冲环境(Peccerillo, 1998; Xia, 2014)。在后碰撞阶段形成的镁铁质岩石,由于继承古板片俯冲改造的地幔源区而通常具有类似于岛弧玄武岩的地球化学组成(El Hassan et al., 2010; Wang et al., 2013)。通过结合前人对南秦岭地区的研究,本文认为铜厂湾辉绿岩应形成于后碰撞阶段,并非岛弧环境下的产物,这一观点对于全面认识勉略宁地区构造成矿过程具有重要意义。三叠纪末期—侏罗纪早期,秦岭造山带的构造体制转换引发了区域一系列岩浆-热液成矿活动。张红等(2015)获得了南秦岭镇安地区钼矿的早侏罗世年龄,同时还获得了围岩花岗岩的早侏罗世年龄,认为该地区成岩成矿时代均位于秦岭碰撞造山之后,是由于造山带垮塌引起的岩浆-热液成矿;齐金忠等(2005; 2008)通过对南秦岭南侧阳山金矿区成矿模式进行总结,认为190 Ma左右阳山矿区内也出现了大规模中酸性岩浆侵位,发育了与成矿相关的岩浆热液活动,这一观点也得到成矿相关花岗岩的蚀变年龄(190 Ma)的支持(杨荣生等, 2006);王瑞廷等(2009)认为,虽然勉略构造混杂带内煎茶岭地区超基性岩形成于晋宁期,但其金矿的形成时代主要为印支期—燕山期,对应于区域上的后碰撞造山过程,为燕山伸展体制下以剪切走滑为主的构造环境。上述事例结合本文研究结果表明,南秦岭构造带存在早侏罗世挤压向伸展过渡体制下形成的岩浆岩,它们指示出早侏罗世区域内存在明显的岩浆活动,因此也可能具有形成热液金矿的可能性,对后续找矿工作具有很好的指示意义。

5 结 论

(1) 陕西勉略宁三角区中铜厂湾金矿区内辉绿岩脉的LA-ICP-MS锆石U-Pb年龄为 (199.6 ± 1.8) Ma,成岩时代为早侏罗世。

(2) 元素地球化学特征显示辉绿岩样品呈亚碱性,属拉斑玄武岩系列,微量元素以亏损Nb、Ta和富集Ba、K等元素为特征,具有类似大陆边缘玄武岩的性质,岩浆源区受到了壳源物质对岩石圈地幔的交代。

(3) 根据岩石特征,并结合区域构造演化历史,认为铜厂湾金矿区内辉绿岩虽然具有类似弧岩浆岩的特征,但它代表了后碰撞阶段减压增温的特殊构造体制下岩浆活动的产物,岩石形成于造山带挤压向伸展过渡阶段。

致 谢 中国地质大学(北京)地球科学与资源学院侯立鸣硕士在锆石测年方面给予了帮助;陕西略阳铀厂沟金矿工程师王立新在野外工作中的建议、帮助并给予的方便,在此一并表示感谢。

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